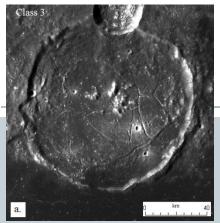
Lunar Floor-Fractured Craters:
Sill Emplacement Models

B. Viscously Relaxed Complex Crater



Lauren Jozwiak James Head ORCH

Th

Short -Wavelength Topography Preserved

Long- Wavelength Topography Amplitude Decreases

C. Complex Crater Intruded by Sill

Floor and Central Peaks
Uplifted

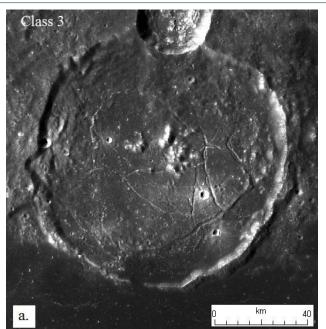
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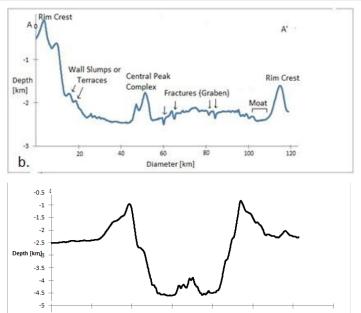
Rim Crest Height
Unchanged

Forms Sill at
Density Boundary

Department of Geological Sciences, Brown University Providence, RI, USA

> NLSI Lunar Science Forum NASA Ames July 18, 2012

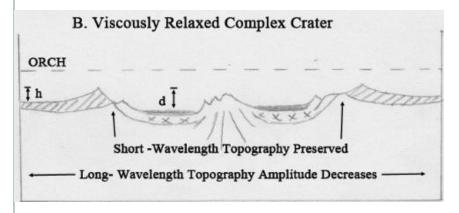


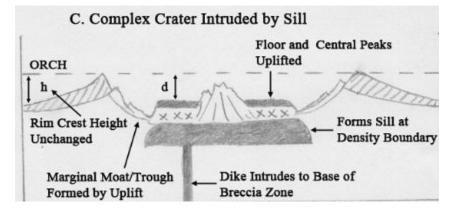


Floor-Fractured Crater

Fresh Crater

Modes of Emplacement





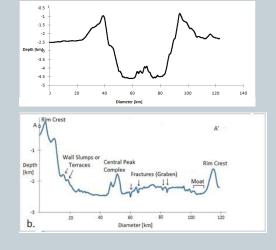
Floor-Fractured Craters (FFCs)

Anomalously shallow craters with fractured floors

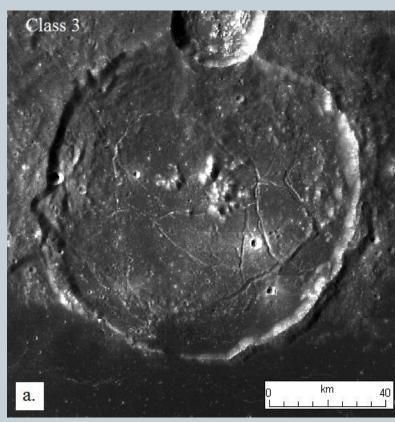
[Schultz, 1976].

Other Characteristics

- Floor moats
- Ridges
- Mare patches
- Dark halos

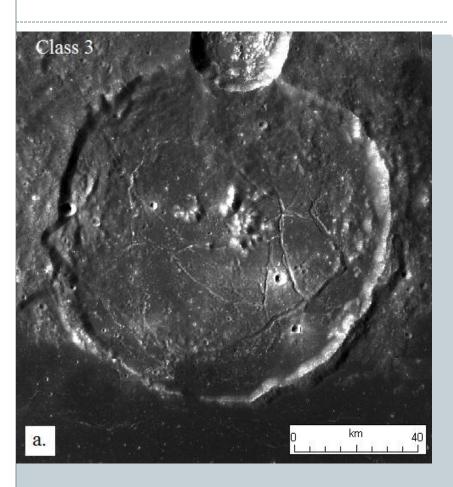


 Characteristics define the morphologic classes.

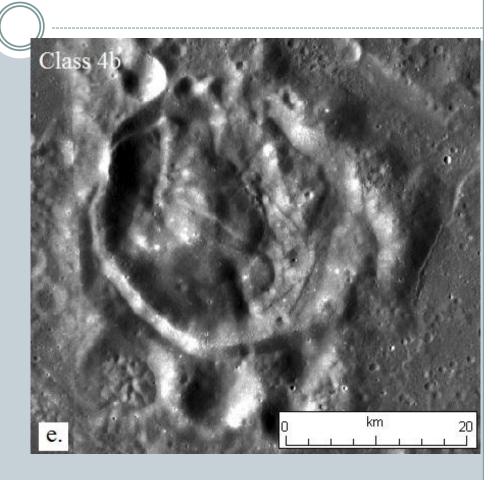


Crater Gassendi, LROC-WAC

Examples of FFC Classes

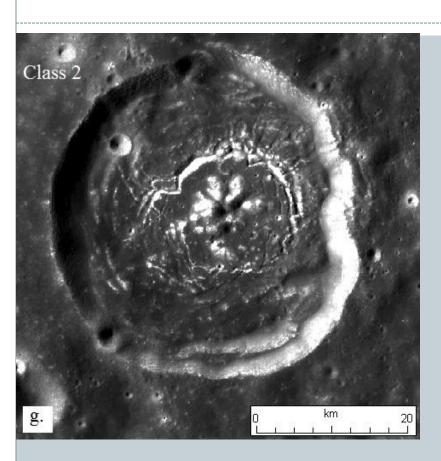


Crater Gassendi Class 3, Wide Moat

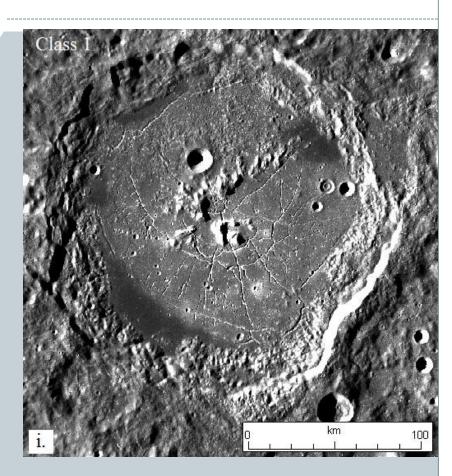


Crater Gaudibert Class 4b, High Ridge V Profile

Examples of FFC Classes

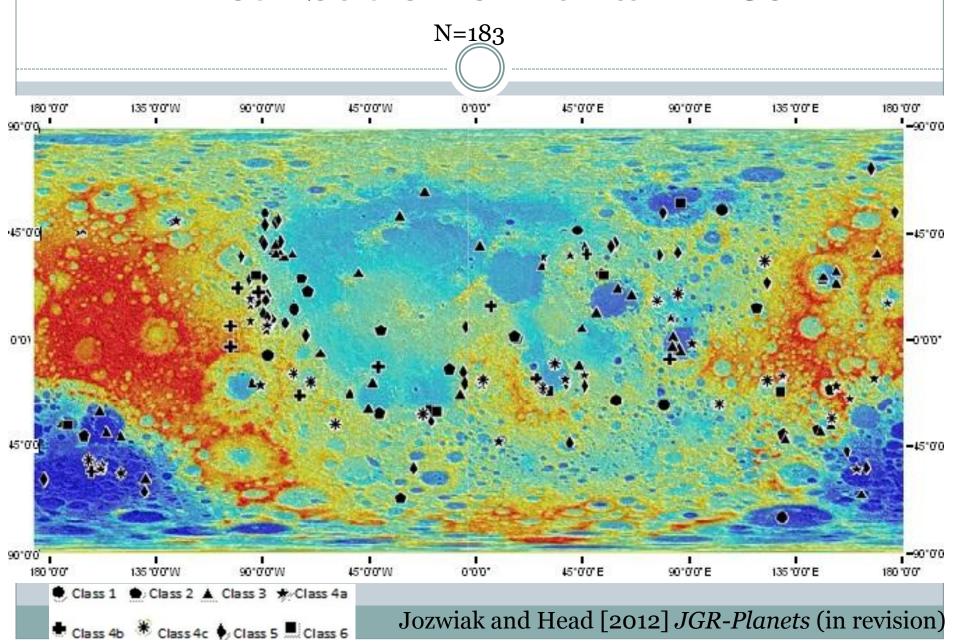


Crater Vitello Class 2, Concentric Central Uplift



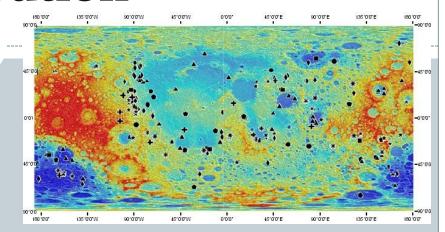
Crater Humboldt Class 1, Mare Patches

Distribution of Lunar FFCs



Distribution

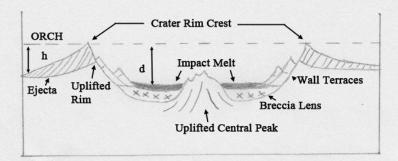
Using LOLA and LROC data,
 classified all FFCs and plotted
 the areal distribution.
 [Jozwiak and Head, 2012, In Revision]



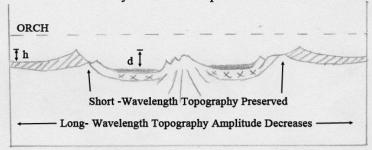
- •Observed relationship between FFC class and its areal distribution.
 - Craters close to basin edges typically have flatter floors, more uplift, and more fractures.
 - Craters farther in the highlands have more convex up floors, less overall uplift.
 - Could be due to 1) thermal effects close to impact basins
 - 2) intrusion effects close to maria and source of magma

Testing Proposed Formation Mechanisms

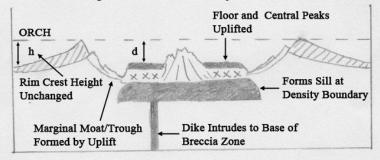
A. Fresh Complex Crater



B. Viscously Relaxed Complex Crater



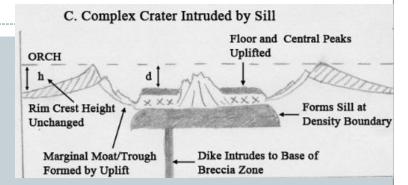
C. Complex Crater Intruded by Sill



- Two proposed formation mechanisms.
- LOLA and LROC have allowed thorough investigation.
- Magmatic Intrusion most probable formation mechanism, supported by:
 - •Significant decrease in floor depth
 - Unchanged Rim Crest Height
 - Lack of crater symmetry
 - Moat features
 - Location far from basin edges
 - •Significant population of small craters

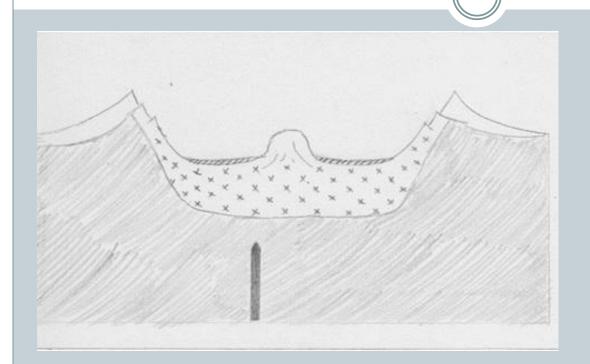
Testing the Mechanics of Magmatic Intrusion

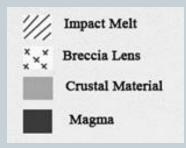
1)Dike propagates from the mantle, driven by a certain pressure.



- 2) Dike stalls at a density barrier caused by the brecciated lens beneath the crater.
- 3) Dike propagates laterally, forming a sill beneath the brecciated lens.
- 4a) Sill inflates, forming a laccolith, and bowing the overlying crater floor.
- 4b) If the yield stress is exceeded at the edges of the laccolith, faulting occurs and uplifts the entire crater floor.

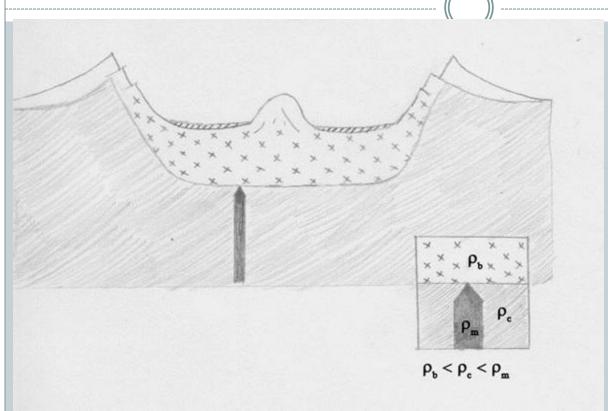
Dike Propagation-Step 1

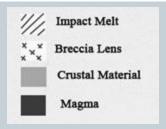




Dike propagates from the mantle with a driving pressure equal to the total magma pressure minus the lithospheric pressure.

Breccia Lens Boundary-Step 2

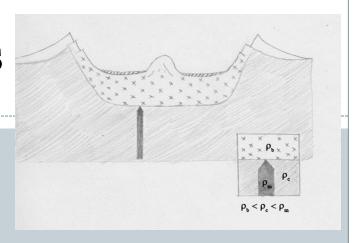




 $ho_{m} = 3300 \text{ kg/m3}
ho_{b} = 2750 \text{ kg/m3}
ho_{c} = 2900 \text{ kg/m3}
ho_{c} = 2900 \text{ kg/m3}
ho_{c}$ [Huang and Wieczorek, 2011]

Breccia Lens

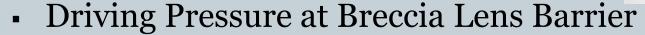
$$K^{+a} = P_d a^{1/2} + [(\pi^{-1} + 0.25)*g*(\Delta \rho)]a^{3/2}$$



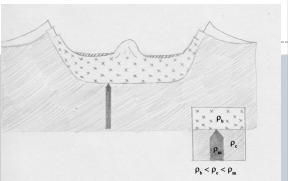
- For a given height, *a*, above the dike origin, a driving pressure of *Pd* is required to fracture rock with the fracture toughness *K*.
- $\Delta \rho$ is the density difference between the intruding magma and the host rock.

Breccia Lens and Driving Pressure

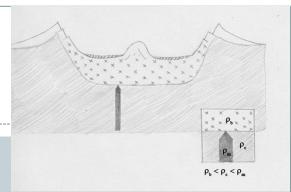
- Driving Pressure to Reach surface
 - 27 MPa through lunar crustal material
 - 32 MPa through breccia lens material

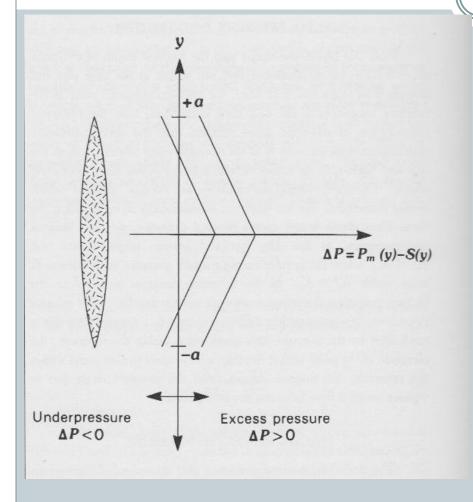


- 21 MPa for intrusion to continue into lunar crustal material
- 29 MPa for intrusion to continue into breccia lens material
- Dike would require an additional 8MPa of driving pressure to continue propagating through the breccia lens.
 - This driving pressure would be high enough to reach the surface if the overlying crater and breccia lens structure were not present
 - Therefore: Crater Floor Breccia Lens Density Barrier is a viable means of halting vertical propagation



Lateral Propagation

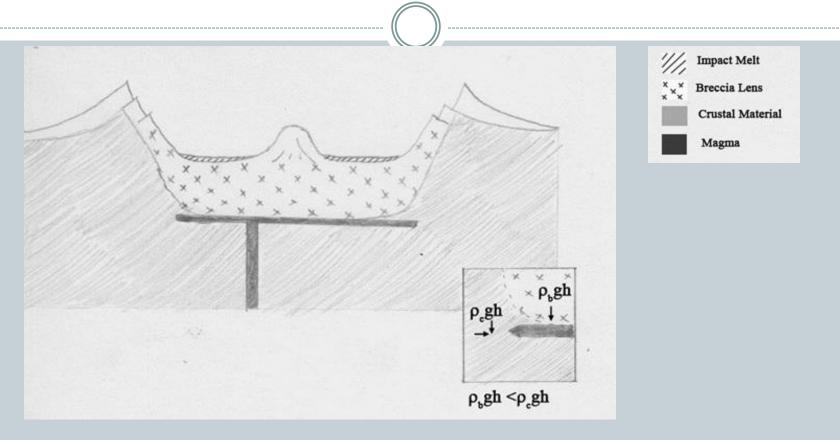




[Rubin and Pollard, 1987]

- Unable to continue vertical propagation, pressure builds below the dike tip.
- This pressure then exceeds the lithostatic pressure, and begins to fracture laterally.
- A lower stress intensity factor on the upper part of the propagating dike ensures that the dike propagates along the base of the breccia lens.

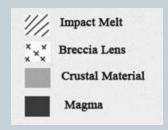
Sill Formation-Step 3



- The sill propagates until it reaches the end of the breccia lens.
- Past the breccia lens, the overburden pressure increases, and the dike encounters a uniform stress state, halting propagation.

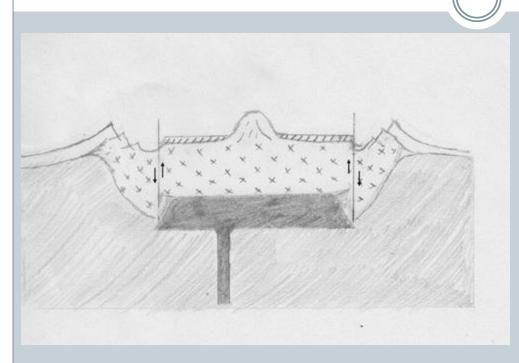
Sill Inflation (Laccolith)- Step 4a

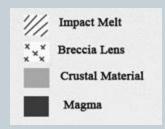




- Magma continues to fill the sill.
- Crystallization of magma at the periphery causes a concentration of magma in the center of the intrusion.
- Extreme flexure in the overlying crater floor, and a convex up floor profile.

Sill Inflation and Faulting-Step 4b





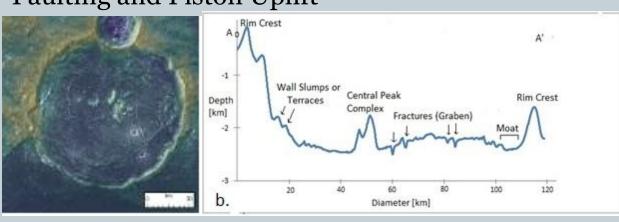
Rapid filling of the sill distributes magma throughout the intrusion volume.

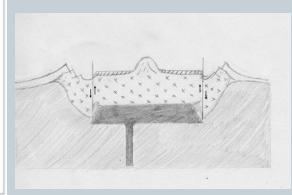
Increased edge stresses overcome the breccia yield stress, and faulting occurs at the periphery.

- Piston-like uplift of the floor replaces flexure with brittle faulting and yields a flatter profile.
- Therefore: Both the Flexural Doming and Piston Uplift are clear consequences of sill intrusion

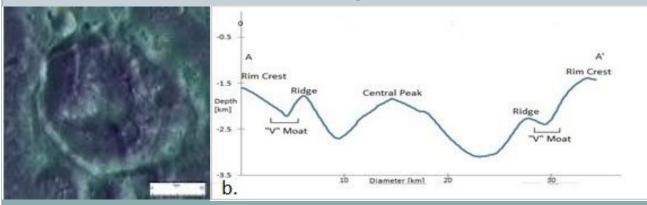
Further Assessing Mode of Origin and Style

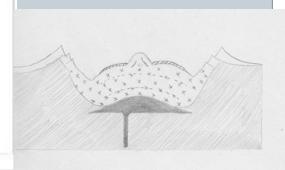
Faulting and Piston Uplift





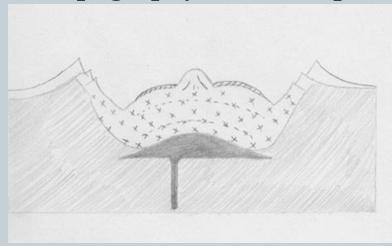
Flexure and Flexural Doming

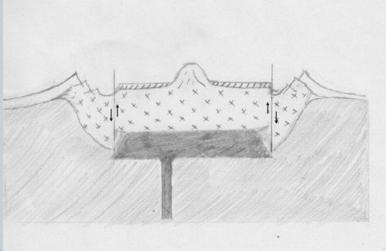


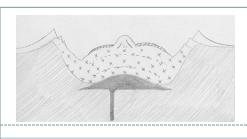


Distinguish by Shape of the Intrusion

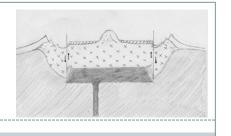
- What do the tails of the intrusion look like?
 - Thin and tapered vs. blunt and snub nosed?
 - Are both present?
 - What determines their shape, and does this affect the crater topography and morhpology?







GRAIL Applications



- Plotting intrusion and crater dimensions to estimate intrusion mass, and associated gravity anomaly, providing predictions to be tested by GRAIL.
- GRAIL can also determine intrusion shape—is the anomaly under a small portion of the crater floor, or the entire floor region? Is this linked to the overlying crater morphology?
- Crater Size Dependence and Intrusion Morphology: look at range of anomaly sizes and shapes with comparison to crater sizes.